RELATIONSHIPS BETWEEN SEA SURFACE TEMPERATURES AND STREAMFLOWS ACROSS THE CORDILLERA DE LOS ANDES

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1. INTRODUCTION

Variations in ocean circulation markedly affect atmospheric conditions in the adjacent marine regions and in the surrounding lands. These changes in the ocean-atmosphere conditions can, in turn, alter the patterns of atmospheric circulation at hemispheric scale, creating a series of anomalies at global scale through teleconnections. Some modes of ocean-atmosphere variability, such as ENSO (El Niño/Southern Oscillation), NAO (North Atlantic Oscillation) and PDO (Pacific Decadal Oscillation), are associated with environmental changes in different regions of our planet. Previous studies in Argentina showed that variations in the spatial patterns of sea surface temperatures (SSTs) affected precipitation fields, and indirectly, the discharge of the Paraná and Uruguary rivers (Mechoso e Iribarren, 1992; Valdés et al., 1999; Berri et al. 1997).

The major goal of this study is to describe the relationships between interannual variations in river discharges along the Andes Cordillera and SST. Basins with different fed-precipitation regimes were selected in order to include streamflow originates from rain (Bermejo river), snowmelt (Atuel river), or a combination of both (Neuquén, Limay and Chubut rivers).

The ocean-atmosphere modes of variability previously mentioned are directly linked to the temporal-spatial evolution of the SSTs and sea level pressure on the adjacent regions. Consequently, in this study we use SSTs as an expression of these modes of variability.

2. MATERIALS AND METHODOLOGY

2.1 Spatial patterns

Monthly streamflow data were provided by the Subsecretaría de Recursos Hídricos, Argentina. The main characteristics of the selected rivers are presented in Table 1. Gridded (on 1° lat. x 1° long.) SSTs for the period 1950 – 2004 were obtained from the Climate Diagnostic Center of the National Ocean and Atmospheric Administration (Reynolds & Smith, 1994).

Oceanic regions were SSTs were statistically linked (or teleconnected) with the temporal variability of the annual mean streamflow (Qma) of the rivers listed in Table 1 were identified. Patterns of spatial correlations between streamflows and SSTs were calculated using the CLIMLAB program (Tanco y Berri, 2000) developed at the International Research Institute (IRI) of the Lamont-Doherty Earth Observatory, University of Columbia (USA). CLIMLAB produces isoline maps connecting SST grids with equal correlation values resulting of comparing streamflows and SSTs.

River	Gage Station	Basin Area (km^2)	Annual Average Discharge (m^3/s)	Data Period
Bermejo	Sarmiento	25.000	354,0	1940-99
Atuel	Angostura	3.800	34,8	1906-99
Neuquén	P. Indios	30.843	308,1	1903-99
Limay	Confluencia	26.688	681,1	1943-99
Chubut	Altares	16.400	46,4	1943-99

Table 1. Gauging stations and characteristics of the analyzed rivers.

2.2 Statistical analysis

In order to establish the statistical significance of the relationships between streamflow and SST, the degrees of freedom in the correlation analyses were adjusted taking into account the autocorrelations in the temporal series. In the process of estimating the statistical significance of the correlation coefficients, the actual number of observations was replaced by the effective number observations (Neff). Neff is calculated as follows (modified from Slonosky et al., 2000).

$$N\text{eff} = N \times (1 - |r_1 r_2|)/(1 + |r_1 r_2|)$$

Where N is the total number of observations and, r_1 and r_2 are the lag-one serial correlation coefficients of the two time series under consideration (Qma and SST). The original equation in Slonosky did not use the absolute value of the correlation product ($r1.r2$). If such product is negative, Slonosky et al (2000) equation might produce values of Neff larger than N.

To determine if the correlation coefficients between discharges and SST are statistically significant we have applied a test based on the distribution of Student.

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Standard Error of $r = \sqrt{(1-r^2)/(N\text{eff}-2)}$

If the r calculated is greater than the tabulated value, which results from multiply the standard error by the t-Student with Neff-2 degrees of freedom, the correlation coefficient is significant.

The Mann-Whitney test of mean population differences was applied to check temporal changes in streamflow across the Andes as well as for SSTs. River and SST time series were split in two 25-year subperiods 1950-1974 and 1975-1999 periods. For the Patagonian rivers (Limay and Chubut), a visual inspection of the records indicates a marked change in streamflows after 1983. In consequence, for the Patagonian rivers the subperiods 1950-1982 and 1983-1999 were considered.

3. RESULTS

3.1 Spatial patterns

Figures 1 to 5 show the spatial correlation patterns between *Qma* and SST for those sectors of oceans where correlations are particularly significant. The *Qma* was calculated taken into account the hydrologic cycle of each river. For the Bermejo river the hydrologic cycle goes from September to August, for the Atuel river from July to June, and for the Neuquén, Limay and Chubut rivers from April to March.
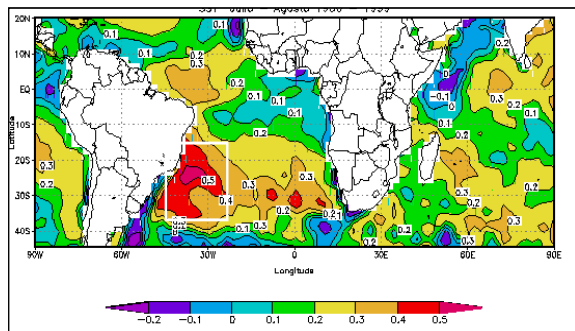

Figure 1.1 Spatial correlation pattern between the *Qma* from the Bermejo river and July-August SSTs for the interval 1950-1999. The white rectangle represents the area over the sea more strongly correlated with the streamflow.

In all figures, positive correlation coefficients between the *Qma* and SST indicate that positive departures in SSTs are related to above long-term mean streamflows. Conversely, negative correlation coefficients suggest that abundant discharges occur in conjunction with negative departures in SSTs.

The Figure 1.1 shows the spatial correlation pattern between the *Qma* from the Bermejo river and July-August SST during the period 1950-1999. The region located in the South Atlantic Ocean off the coast of South America between 16º and 36° S is significant correlated with the *Qma* from the Bermejo

river. For the sector located between 20° and 25° S, r reaches values above 0.5. During November-December, the region of highest correlations indicated in Figure 1.2 maintains correlation values superior to 0.3, however, the ocean sector more strongly correlated with the streamflow appears now above the Indian Ocean adjacent to South Africa. This region presents correlation values in the order of $r = 0.55$.
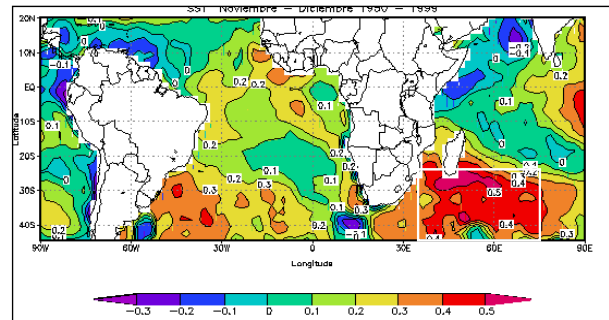

Figure 1.2 Spatial correlation pattern between the *Qma* from the Bermejo river and November-December SSTs during the interval 1950-1999.
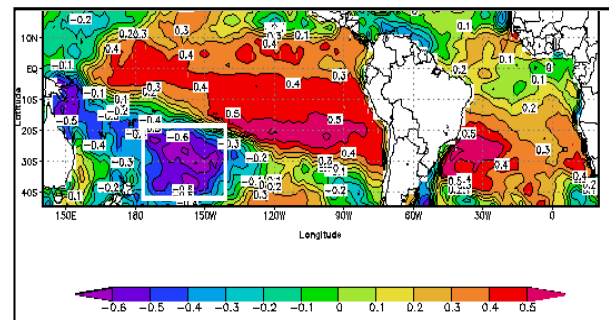

Figure 2.1 Spatial correlation pattern between the *Qma* from the Atuel river and July-August SSTs for the interval 1950-1999. The are on the Pacific next to the Australian coast (22° and 40° S), presents correlations with the Atuel streamflow in the order of $r = 0.7$.
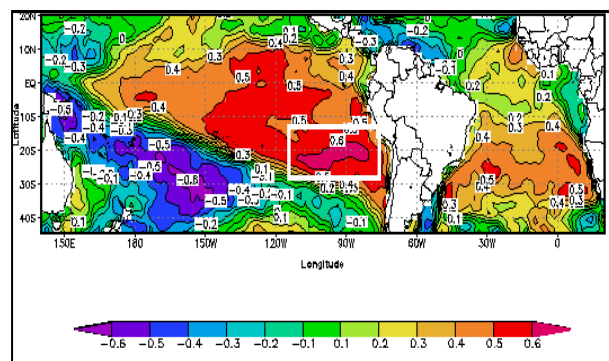

Figure 2.2 Spatial correlation pattern between the *Qma* from the Atuel river and September-October SSTs for the interval 1950-1999. The Pacific area off to the South American coast (10°-30° S), presents correlations in the order of $r = 0.6$.
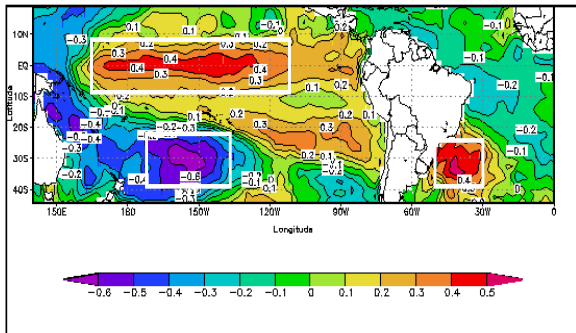
Figure 3 Spatial correlation pattern between the *Qma* from the Neuquén river and June-July SSTs during the period 1950 – 1999. White boxes delimit the areas on the sea more strongly correlated with the streamflow. The three oceanic regions strongly related to the Neuquén river streamflow are the tropical Pacific and the western subtropical Pacific in a positive way, and the western Subtropical Pacific in a negative way.
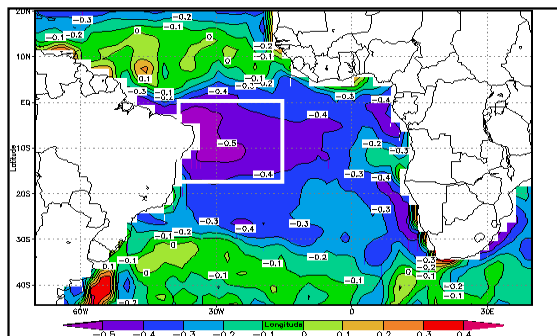

Figure 4.1 Spatial correlation pattern between the *Qma* from the Limay river and January-February SSTs during the interval 1950-1999. Correlation coefficients is in the order of $r = -0.5$ are recorded in the area of the Atlantic Ocean in front of the NE Brazilian between 0° and 20° S.
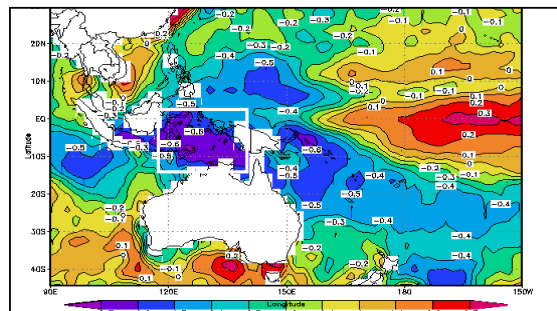

Figure 4.2 Spatial correlation pattern between the *Qma* from the Limay river and October-November SSTs during the interval 1950-1999. The region located in the west Pacific to the north of Australian reaches correlation coefficients about $r = -0.7$.
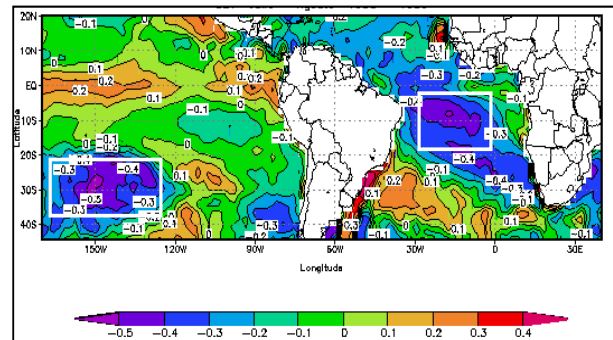

Figure 5. Spatial correlation patterns between the *Qma* from the Chubut river and July-August SSTs during the period 1950-1999. The regions located on the southwestern Pacific Ocean (20º - 40º S) and in the tropical Atlantic (0º - 20º S) are significant correlated with the Chubut river streamflow ($r = -0.5$ and $r = -0.4$, respectively).

Figures 2.1 and 2.2 show the spatial correlation patterns between the annual discharges of the Atuel river and SSTs for July-August and September-October, respectively. In both patterns the east and southwest regions of the Tropical-Subtropical Pacific Ocean between 10° - 30° S and 22°-40° S appear strongly correlated with the Actual river streamflow. Correlation coefficients reach absolute values between $r = 0.6$ and 0.7.

Figure 3 represents the spatial correlation pattern between the Neuquén river streamflow and June-July SSTs. Correlation values are in the order of $r = -0.6$ in the South Pacific east of the Australian coast, $r = 0.5$ in the South Atlantic sector close to Santa Catarina (Brazil), and $r = 0.4$ in the Tropical Pacific.

Figures 4.1 and 4.2 represent the spatial correlation patterns between the *Qma* from the Limay river and January-February, October-November SSTs, respectively. For January-February, SSTs in the tropical Atlantic off the Brazilian coast show correlation coefficients around $r = -0.5$, whereas values close to $r = -0.7$ are recorded for the western Pacific region during October and November.

Finally, the Figure 5 shows the spatial correlation patterns between the *Qma* from the Chubut river and SST across the Pacific and Atlantic Oceans. Two oceanic regions located in the Subtropical Pacific between the 20° and 45° S and in the Tropical Atlantic between 0° - 20° S, are inversely correlated with the Chubut river streamflow.

3.2 Statistical analysis

Table 2 lists the correlation coefficients between the *Qma* from the five selected rivers and SSTs from the areas indicated with white rectangles in Figures 1 to 5. Correlation coefficients between streamflow and SST are indicated as calculated r_c, while the critical correlation value based on *N*eff is designed as tabulated r_t. For all comparisons, the calculated r_c is higher than the tabulated r_t. The level of significance provides a measurement of the strength of the relationships.

For most of the analyzed rivers the Mann-Whitney test indicates that means for the selected sub-periods are significantly different, with an increase in mean discharge

during the period 1975-1999 in the basins of the Bermejo (highly significant) and Atuel (statistically significant) rivers, whereas a reduction was recorded for the Limay and Chubut rivers. There are no significant changes in the Neuquén river between these periods (Table 3).

Significant results were obtained using the Mann-Whitney test in the process of detecting changes in SSTs across the regions significantly correlated with the river streamflows across the Andes. In general higher SSTs during the interval 1975-1999 have been recorded in the Pacific, Atlantic and Indian Oceans, south of the Equator. Table 3 provides the Mann-Whitney test results applied to SSTs during different intervals. Changes detected in streamflow regimes (increase or decrease) and in SSTs (warming or cooling) are indicated in Table 4.

River	Region of the Ocean	N	Neff	r_c	r_t	α(%)	Figure
Bermejo	Atlantic W	50	40	0.51	0.44	99.9	1.1
Bermejo	Indian W	50	40	0.61	0.41	99.9	1.2
Atuel	Pacific SW	50	40	-0.7	0.36	99.9	2.1
Atuel	Pacific SE	50	33	0.6	0.42	99.9	2.2
Neuquén	Pacific SW	50	48	-0.7	0.37	99.9	3
Neuquén	Atlantic SW	50	49	0.52	0.44	99.9	3
Neuquén	Tropical Pacific	50	49	0.39	0..36	99.0	3
Limay	Atlantic W	50	44	-0.5	0.44	99.9	4.1
Limay	Pacific W	50	48	-0.7	0.36	99.9	4.2
Chubut	Pacific S	50	46	-0.5	0.44	99.9	5
Chubut	Tropical Atlantic	50	39	-0.4	0.39	99.5	5

Table 2. Correlation coefficient between the SST from some selected sea areas and streamflows from the Andean rivers (r_c: correlation coefficient calculated, r_t: correlation coefficient tabulated, α: Level of Significance).

River	α(%)	Class.	Region of the Ocean identified	α(%)	Class.
Bermejo	99.9	HS	Atlantic W	99.9	HS
Bermejo			Indian W	99.9	HS
Atuel	99.0	SD	Pacific SE	99.9	HS
Atuel			Pacific SW	95.0	NSD
Neuquén	95.0	NSD	Pacific SW	95.0	NSD
Neuquén			Atlantic SW	95.0	NSD
Neuquén			Tropical Pacific	95.0	NSD
Limay	99.0	SD	Atlantic W	99.0	SD
Limay			Pacific W	95.0	MD
Chubut	99.0	SD	Pacific S	95.0	NSD
Chubut			Tropical Atlantic	99.9	HS

Table 3. Mann-Whitney Test for differences in mean populations between subperiods. Difference between intervals are classified as: HS: highly significantly; SD: significantly different; MD: marginally different; NSD not significantly different.

4. CONCLUSIONS

As part of an integral project intended to analyze the impact of climatic changes on water resources across the Argentine Andean Cordillera, we identified in this study the ocean regions whose changes in temperature are strongly correlated with Andean river discharges. The influence of changes in SST over the climate of adjacent regions, and remote regions through teleconnections, has been largely acknowledged and our study provides additional evidence of these phenomena for the austral region of South America.

River	Observed Change	Region of the Ocean	Observed Change
Bermejo (1)	Increase	Atlantic W	Warming
Bermejo (1)		Indian W	Warming
Atuel (1)	Weak Increase	Pacific SE	Warming
Atuel (1)		Pacific SW	Weak Warming
Neuquén (1)	Weak Decrease	Pacific SW	Not change
Neuquén (1)		Atlantic SW	Weak Warming
Neuquén (1)		Tropical Pacific	Weak Warming
Limay (2)	Decrease	Atlantic W	Warming
Limay (2)		Pacific W	Warming
Chubut (2)	Decrease	Pacific S	Not change
Chubut (2)		Tropical Atlantic	Warming

Table 4. Changes observed in mean streamflow and SST anomalies between the periods 1950-1974 and 1975-1999 (1), and between the periods 1950-1982 and 1983-1999 (2).

The spatial correlation patterns show geographical differences among the oceanic regions that influence changes in the discharges of rivers located in different sectors of the Andes: northwestern Argentina, the Central Andes of San Juan and Mendoza, and Northern Patagonia. The streamflow of Bermejo river, located north of 24ºS, is strongly associated with temperatures in the subtropical Atlantic off the South American coast. These results are consistent with previous studies pointing at the subtropical Atlantic as one of the main sources of humidity for Northwestern Argentina (Barros et al. 2000; Vuille et al. 2004).

In response to the topography of the oceanic basins and the associated oceanic and atmospheric circulations, there are ocean regions where temperature evolves in a similar way. Recently, it has been observed that temperatures in the Subtropical Atlantic Ocean co-vary in a similar way with temperatures in the Subtropical Indian Ocean close to the African coast (Fauchereau et al., 2003). This can explain the significant correlations between the Bermejo river streamflow and the Subtropical Indian Ocean during the November-December period.

Relationships between precipitation in the Central Andes and Tropical Pacific SSTs have been observed by several researchers (Montecinos and Aceituno, 2003). In agreement with these observations, we found a strong relationship between variations in SST on the Tropical Pacific and the Atuel river streamflow. Relationships between SSTs and the discharge of rivers in the northern Patagonia are more difficult to interpret. It is quite probable that many of the observed spatial correlation patterns are owed to the way in which SSTs co-vary among different ocean regions. On the other hand, preliminary studies indicate that the Puelo river streamflow, on the western slope of the Patagonian Andes, is strongly regulated by changes in atmospheric circulation at high altitudes (Lara et al., submitted). In particular, a strong relationship between the Antarctic Oscillation (Gong, 1999) and the summer-fall Puelo river discharge has been observed. Unfortunately, our SST data set does not included records south of 45° S, hampering to identify the influence of high-latitude ocean circulation on the northern Patagonian (Limay and Chubut) river discharges.

In the mid 70's, a significant change was observed in the hydrologic regime of most river across the Andes with larger discharges for the rivers located in northwestern Argentina (Bermejo) but a reduction for the Patagonian (Limay and Chubut) rivers. Similar changes in SST anomalies were recorded in oceanic regions teleconnected with streamflow variations in many rivers along the Andes (Table 4). Some of the changes in streamflow during the past 50 years are related to warmer SST on the Pacific, Atlantic and Indian Oceans south of the Equator.

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